

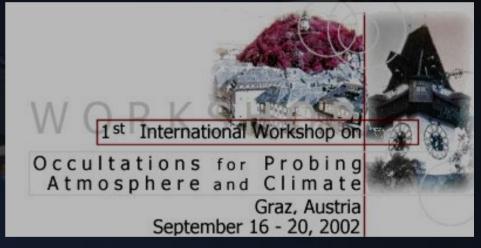
IGAM/UG - Institute for Geophysics, Astrophysics, and Meteorology / University of Graz

Atmospheric Remote Sensing and Climate System Research Group

ARSCliSys – on the art of understanding the climate system







Stratospheric Temperature and Ozone Sounding with ENVISAT/GOMOS Stellar Occultation

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Middle Atmospheric Ozone Sounding by the ENVISAT/GOMOS Stellar Occultation Sensor



Outline

ENVISAT/GOMOS Stellar Occultation Sensor

Forward Modeling and Retrieval

First Results

Summary and Conclusions





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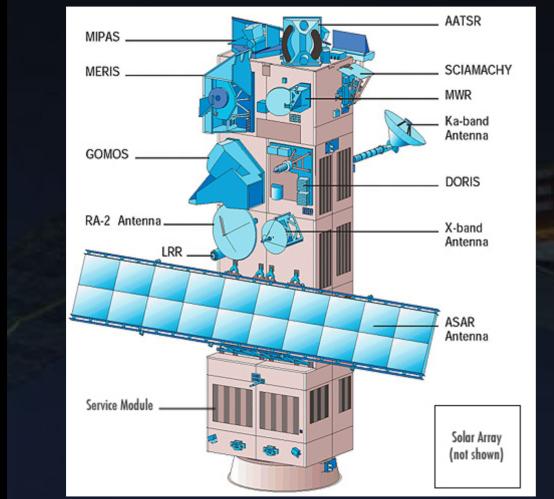


the ENVISAT mission

- Launch with Ariane 5 on 1st March 2002, 1:07:59 UT
- Mission duration > 5 years
- Sun-synchronous orbit: inclination 98 deg, ~800 km altitude
- Ten instruments on-board: ASAR, GOMOS, LRR, MIPAS, MERIS, MWR, RA-2, AATSR, DORIS, SCIAMACHY
- Simultaneous monitoring of land, oceans, ice fields and the atmosphere













the GOMOS sensor concept

- Atmospheric profiles by stellar occultation method (~15 90 km)
- Two Spectrometers
 - A: 248 371 nm (UV), 387 693 nm (VIS)
 - B: 750 776 nm (IR1), 915 956 nm (IR2)
- Two broad-band Photometers with 1 kHz sampling rate
 - 466 528 nm, 644 705 nm
- Self-calibrated transmission data (normalized intensities)
- Profiles of O₃, NO₂, NO₃, BrO, OClO, O₂, water vapor, air density, temperature and turbulence





the star tracking unit

- SFM (steering front mechanism) acquires the star image
 - angular range: ±26°
 - maintains the position and rotates the mirror
- SFA (steering front assembly) with 5 Hz sampling
 - azimuth angle: -11° to 91°
 - elevation angle: 61.7° To 69°
- SATU (star acquisition and tracking unit) with 100Hz sampling
 - nominal and redundant star tracker
 - image of the observed star at focal length of 630 nm
 - SATU errors are between ~ ±10µrad in both directions



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the forward model

• Observed-Transmission model basic formula (no defocusing):

$$T_{obs} = \int_{\Delta t} \int_{\Delta \theta \Delta \lambda} T_{\nu}(\lambda, \theta, t) W(\lambda, \theta, t) d\lambda d\theta dt$$

Modeled atmospheric transmission given by Beer-Bouguer-Lambert law

$$T_{\nu}(t) = \frac{I_{\nu}(t)}{I_{\nu}(0)} = \exp\left[-\int_{x_{S}(t)}^{x_{G}(t)} \sum_{i} n_{i}(s')\sigma_{i\nu}(s')ds'\right]$$

- no finite field of view integration needed, no aerosol extinction used
- CIRA-86 or MSIS-90 models used for bulk atmosphere and standard profiles for trace species O₃, NO₂, and NO₃





the forward model

- Absorption cross sections for O₃, NO₂, and NO₃
- Scattering cross section (Rayleigh scattering)
- realistic refractive path by fast 3D raytracing
- realistic geometry used and Earth's shape approximated by WGS-84
- channel selection (260, 280, 290, 295, 305, 310, 320, 328, 330, 600, 610 nm)
 - O₃ Hartley band (~200 300 nm)
 - slight temperature dependencies and NO₂ sensitivity
- standard GOMOS 2Hz sampling of transmission data → 1.5 km spacing





connecting the forward and the retrieval model

the forward model reads

$$\mathbf{y} = \mathbf{K}(\mathbf{x}) + \mathbf{\varepsilon}$$

- y, x... measurement and state vector
- K... forward model operator, Jacobian matrix (m×n)
- ε... measurement error vector
- rows of Jacobian K can be interpreted as "weighting functions"
- the direct inversion reads

$$\mathbf{x}_r = \mathbf{K}^{-g} \mathbf{y}$$

- ill-conditioned problem at high altitudes
- over-determined for m>n





the retrieval model

- Optimal estimation
 - incorporates sensibly a priori knowledge
 - statistically optimal combination of unbiased measurements and prior data
- fast converging iterative optimal estimation scheme

$$\mathbf{x}_{i+1} = \mathbf{x}_{ap} + \mathbf{S}_i \mathbf{K}_i^T \mathbf{S}_{\varepsilon}^{-1} \left[(\mathbf{y} - \mathbf{y}_i) + \mathbf{K}_i (\mathbf{x}_i - \mathbf{x}_{ap}) \right]$$

$$\mathbf{S}_{i} = \left(\mathbf{K}_{i}^{T} \mathbf{S}_{\varepsilon}^{-1} \mathbf{K}_{i} + \mathbf{S}_{ap}^{-1}\right)^{-1}$$

- **S**_e... observation and forward modeling error covariance matrix
- **S**_r... retrieval error covariance matrix
- **S**_{ap}... a priori error covariance matrix
- **x**_{ap}... a priori profile
- **x**_{i+1}... retrieved profile (iteration i)





bending angles and refractive index

- exploitation of bending angles (α) data (SFA/SATU)
- direct inversion of the refractive index (Abel Transform)

$$\alpha(a) = 2a \int_{r=r_0}^{r=\infty} \frac{1}{\sqrt{n^2 r^2 - a_0^2}} \frac{d \ln(n)}{dr} dr$$

$$n(r_0) = \exp\left[\frac{1}{\pi} \int_{\alpha = \alpha(a_0)}^{\alpha_0} \ln\left(\frac{a(\alpha)}{a_0} + \sqrt{\left(\frac{a(\alpha)}{a_0}\right) - 1}\right) d\alpha\right]$$

- r... radius at tangent point
- a... impact parameter





optimized bending angles

• statistical optimization to find the most probable bending angle profile

$$\alpha_{opt} = \alpha_b + (\mathbf{B}^{-1} + \mathbf{O}^{-1})^{-1} \mathbf{O}^{-1} (\alpha_o - \alpha_b)$$

- α_{opt} ... most probable bending angle profile
- α_o ... observed bending angle profile (ionosphere-corrected)
- α_b ... background (*a priori*) bending angle profile (from climatology)
- O observation error covariance matrix
- **B** background error covariance matrix
- high altitudes: α_{opt} determined by climatology
- \bullet low altitudes: $\alpha_{\it opt}$ determined by observed data





refractivity, pressure and temperature

 refractivity (N) converted to pressure (p) by integration of the hydrostatic equation

$$N(r) = (n(r) - 1) \times 10^6$$

$$p(r) = -b_1 \int_z^\infty N(r')g(r')dr'$$
 with $b_1 = 4.489 \times 10^{-5} \text{ hPa s}^2/\text{m}^2$

• temperature obtained from p and N by equation of state

$$T(r) = c_1 \frac{p(r)}{N(r)}$$
 with $c_1 = 82.86$ K/hPa at 300 nm slightly λ -dependent



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error modeling, first ten O₃ a priori error patterns

• error patterns from **S**_{ap} (20% uncertainty, 6 km vertical correlation length)

$$\mathbf{S}_{ap} = \sum_{i} \mathbf{e}_{i} \mathbf{e}_{i}^{T}$$
 with $\mathbf{e}_{i} = \sqrt{\lambda_{i}} \mathbf{l}_{i}$

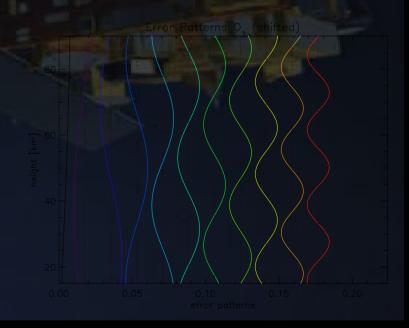
$$\mathbf{e}_i = \sqrt{\lambda_i} \mathbf{l}_i$$

• consistent **x**_{ap} based on "true" profile

$$\mathbf{x}_{ap} = \mathbf{x}_{true} + \Delta \mathbf{x}$$

$$\Delta \mathbf{x} = \sum_{i} a_{i} \mathbf{e}_{i}$$

■ a_i random deviates drawn from a normalized Gaussian distribution



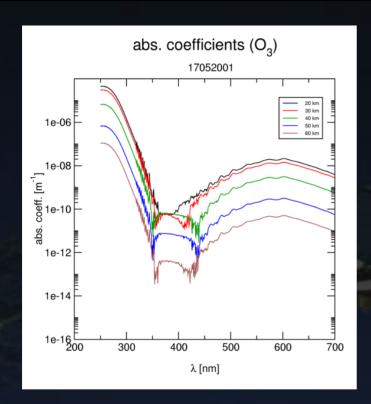




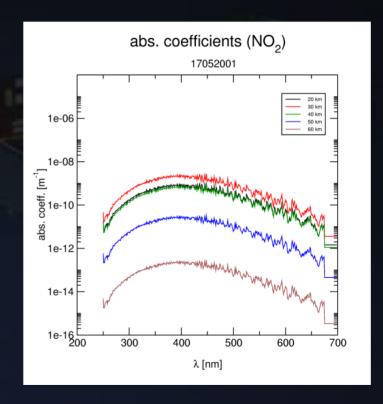






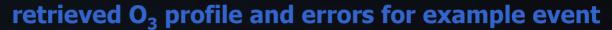


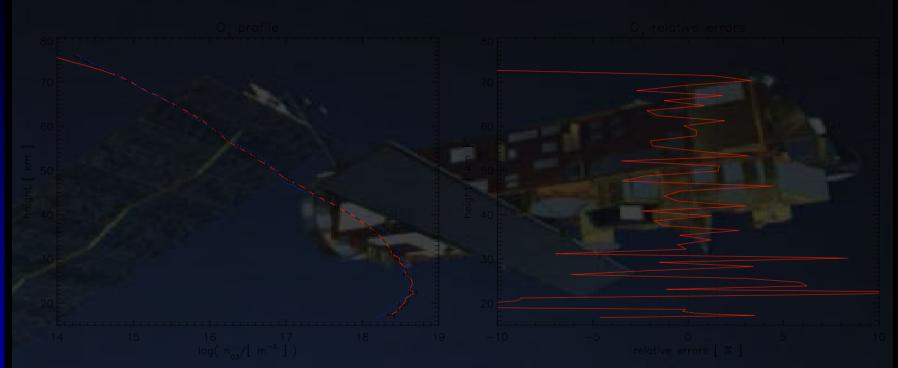
absorption coefficients of O₃ and NO₂











- transmission errors: GOMOS-type values of $0.01\times(T)^{-1/2}$
- no correlation between channels assumed

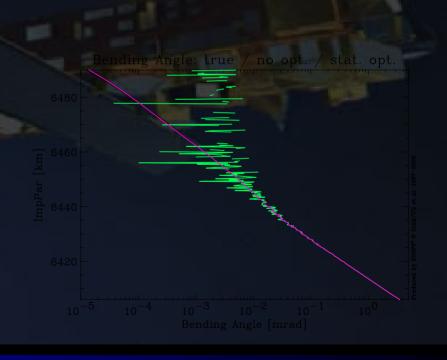


statistically optimized bending angles

• analytical background (a priori) covariance matrix

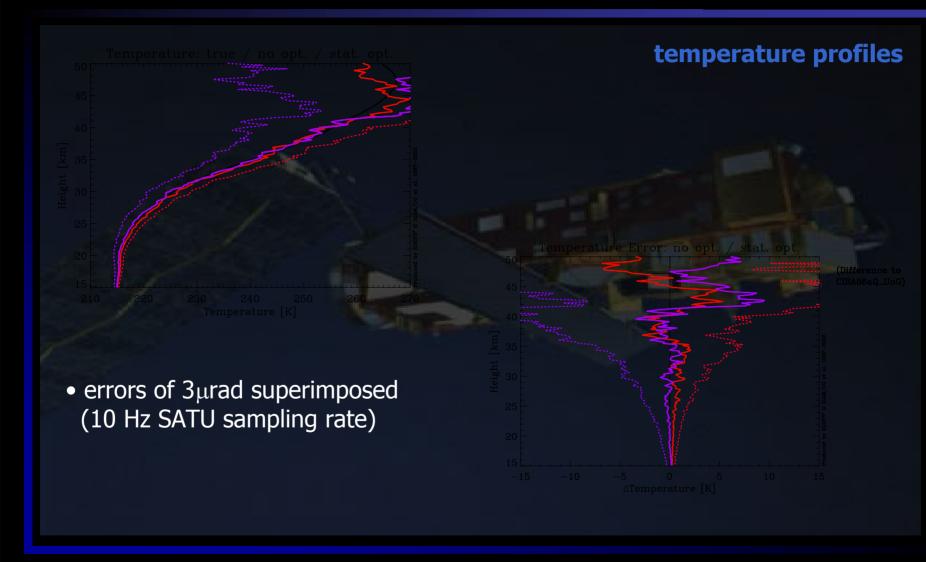
$$B_{ij} = \sigma_i \sigma_j e^{-\left[\left|a_i - a_j\right|/L\right]}$$

- a_i, a_j... *i* th and *j* th impact parameter values
- L... error correlation length = 6 km
- *σ_{i,j}* ... 20% uncertainty











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Summary and Conclusions



the forward model

- realistic observed transmissions from 0 120 km
- channel selection within 260 340 nm (channels outside possible)
- raytracing to be enhanced by GOMOS-measured bending angles
- along-ray integration to be updated by including aerosol extinction



Summary and Conclusions



the retrieval model

- simultaneous O₃ and NO₂ retrieval
- errors of retrieved O₃ profiles within expected ranges (<2%)
- performance to be further improved by
 - channel clustering
 - simultaneous estimation of O₃, NO₂, refractivity, and temperature
- best performance between 15 and 35 km (<1 K below 25 km, < 2 K
- below 35 km) with statistical optimization
- temperature retrieval between 15 and 50 km
- errors > 3 K above 30 km without statistical optimization
- further algorithm enhancements for use with
 - real level1b SFA/SATU data



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